

# **Mongolian Geoscientist**



## Original article

### Stratigraphy of the Khuvsgul Group, Mongolia

Eliel Anttila<sup>1\*</sup> <sup>(D)</sup>, Francis Macdonald<sup>1</sup> <sup>(D)</sup>, Uyanga Bold<sup>2</sup> <sup>(D)</sup>

<sup>1</sup>Department of Earth Science, University of California Santa Barbara, CA, 93106, USA <sup>2</sup>Department of Research and Cooperation, Mongolian University of Science and Technology, Ulaanbaatar 14191, Mongolia

\*Corresponding author: eanttila@ucsb.edu , ORCID: 0000-0002-7563-2710

### ARTICLE INFO

Article history: Received 1 March, 2021 Accepted 20 June, 2021

### ABSTRACT

The Khuvsgul Group (Khuvsgul Province, Mongolia) is a Late Neoproterozoic to Cambrian carbonate-dominated succession that includes minor glacial diamictite and one of the largest known ore-grade phosphate deposits in the world. These strata, which have experienced low-grade metamorphism, are exposed in the Khoridol-Saridag Range on the western margin of Lake Khuvsgul. Since 2017, new geologic mapping and field studies have been conducted in the Khuvsgul region. During the course of this work, it has become necessary to restructure the stratigraphic framework of the Khuvsgul Group in order to better facilitate geologic mapping, stratigraphic observations, and regional correlations. We have divided the lower Khuvsgul Group into four distinct formations spanning the Cryogenian and Ediacaran, each of which encompass strata associated with the Sturtian glaciation, Cryogenian non-glacial interlude, Marinoan glaciation, and basal Ediacaran transgression respectively. The phosphorites of the Khuvsgul Group are now included within a new distinct formation, while the overlying Cambrian carbonates and siliciclastic rocks have been further subdivided to streamline mapping and correlation efforts. The stratigraphic framework outlined below will simplify identification and differentiation of Khuvsgul Group rocks in the field and provide a foundation for the interpretation of Khuvsgul Group strata within the context of the changing climatic, tectonic, and paleoenvironmental conditions of the late Neoproterozoic and early Cambrian.

*Keywords:* sedimentology, phosphorite, Cryogenian, Ediacaran, Cambrian, CAOB.

### **INTRODUCTION**

Located on the northern Tuva-Mongolia terranes (Ilyin, 1971; Kuzmichev, 2015; Bold et al., 2019), and exposed in a north-south belt along the ridgelines above the west shores of Lake Khuvsgul (Figs 1 and 2), the Khuvsgul Basin hosts the late Neoproterozoic to Cambrian Khuvsgul Group (Ilyin and Ratnikova, 1981;

Macdonald and Jones, 2011). The Khuvsgul Group is a ~3 km-thick succession composed predominantly of limestone and dolomite, with minor diamictite, phosphorite, and siliciclastic rocks. Early geological investigation of the Khuvsgul Group centered around the ore-grade Khuvsgul phosphorite (Ilyin, 1973; Ilyin and Ratnikova, 1981). More recent work has focused

<sup>©</sup> The Author(s). 2021 **Open access** This article is distributed under the terms of the Creative Commons Attribution 4.0 International License (https://creativecommons.org/licenses/by/4.0/), which permits unrestricted use, distribution, and reproduction in any medium, provided you give appropriate credit to the original author(s) and source, provide a link to the Creative Commons license, and indicate if changes were made.

on Cryogenian Snowball Earth glaciations recorded by the diamictite units (Macdonald and Jones, 2011) and the late Ediacaran to Cambrian Doushantuo-Pertatanka-type microfossil assemblage within the phosphatic strata (Anderson et al., 2017; 2019). Over the past half-century, workers have developed a variety of stratigraphic schemes for the Khuvsgul Group, with correlations to putatively equivalent units on the Siberian side of the Russian-Mongolian border (Osokin and Tyzhinov, 1998; Kuzmichev et al., 2001; Vishnevskaya and Letnikova, 2013), or the southwest margin of the Zavkhan terrane (Macdonald and Jones. 2011, Smith et al., 2016, Bold et al., 2016a,b). Stratigraphic frameworks proposed by previous workers are summarized in Fig. 3. Here we propose an updated stratigraphic framework geological based on new mapping, lithostratigraphic observations. and

biostratigraphic constraints (see Figs 2, 3, 4, and 5).

Based on the regional basinal stratigraphic divisions outlined by Macdonald and Jones (2011), our new stratigraphy further divides map units on the basis of discrete lithostratigraphic trends observed throughout the Khuvsgul paleobasin. First, we propose that the Ongolog Formation, which includes diamictite and carbonate, be separated into the Ongolog Formation (glacial facies) and Bakh Formation (non-glacial facies). The Bakh Formation is further subdivided, based on basinallyconsistent lithological changes, into the Khurts, Bumbulug, and Salkhitai members. Next, strata formerly incorporated into the Kheseen Formation (Macdonald and Jones, 2011) are subdivided: the newly defined Shar Formation includes a diamictite and breccia sequence, while а cap carbonate and overlying



**Fig. 1.** Location and geologic context of northern Mongolia. A: Geopolitical map of northern Mongolia. B: Tectonic map of the northern Mongolian portion of the Central Asian Orogenic Belt, modified from Bold et al. (2016b, 2019) and Kuzmichev (2015). C: Generalized geologic map of the Khoridol-Saridag Range and Darkhat Valley. A detailed geologic map of the area outlined in red is shown in Fig. 2.



**Fig. 2.** Geologic map of the eastern and central Khoridol-Saridag Range, Khuvsgul Aimag, Mongolia. Locations of measured stratigraphic sections are depicted with lavender lines. Only type-locality measured sections are labeled, as are river drainages or gorges ("gols") relevant to locations discussed in this work. Note that the Shar Formation and Khirvesteg Formation are depicted in this figure as a single unit, as the Khirvisteg Formation is generally too thin to be clearly discernible at the map scale presented here.

			unconformity	diamicute dolomite limestone siliciclastics chert volcanics	ithology, and
	یے 1000 میں 1000 میں	300m -500m	50		reported 1
This Study	irkhelnuur Fm		K Member Mem	Darkhat Group	presentations of 1
	Set. 2-3? Sotom. Set. 2-3?	TechDal. To	Cryogenian	nsinoT	re
	Cambrian		proterozoic	Nec	loc
	10077	1071	1000 00h		arte
		\$			ပ ၂ ပ
2011; 2011; 2019	Erkhelnuur Fm	Khesen Fm	Oudoluk Fm	m∃ nessıA	ure
ones, ones, derso 2017,	dnou	S Ingusdud		Darkhat Group	clat
ΣŢΥ	nsindmeD	Ediac.	nsinapoviD	nsinoT	ene
			Diozorato	ICON	
					ม ม
2003					din
et al,	rkhelnuur Fm	E Seu Fm	Ongolik Fm{ Kh	m 7 neserA	clu
denko	roup (Ukhutologoi	S Ingusdudy		Darkhat Group	E.
Demic	Adtabanian Botomian	neitommoT	nsibn9V	Upper Riphean	ks.
_	neindme2				vor
	ш005—н ш090Z				nev
al, 00;			ИНИНИИ		frar
al, 20 la et a 5	Erkheinuur Fm	MJ uəsəy	m <sup>3</sup> AulopnO	m∃ neserA	с Т
12 5	ADOI OF THE OF T				
allo et rnamja 2015		Khubsanaul G	Faidearai		
Zhegallo et Dornamja: 2015	Adtabanian Botomian Adtabanian Botomian	Tommotian Khubsugul G	Ediacaran	nisandin raddo Tonian-Cryogenian	rapl
Zhegallo et Dornamja: 2015	Cambrian Botomian Botomian	nsitommoT Rhubuqudul R	Vendian Ediacaran	Upper Riphean Tonian-Cryogenian	atierapl
998; Zhegallo et cova, Dornamja: 2015	Cambrian Botomian Botomian	Tommotian	Vendian Ediacaran	nenogo *cd/red neinogo *cd/red	stratigrapl
Letnikova, Dornamja. 2015	Cambrian Botomian Botomian Botomian Cambrian Cam	-2000	-³00m –⊫ −m005 Vendian Ediacaran	ozr	us stratigraph
a Tyzhinov 1998; Zhegallo et i a and Letnikova, Dornamja: 2013	Lower Upper 1000-1350m 1000-1300m 1000-1000m 1000-1000m 1000-1000m 1000-1000m 1000-1000m 1000-1000m 1000-1000m 1000-1000m 1000-1000m 1000-1000m 1000-1000m 1000-1000m 1000-1000m 1000-100000000000000000000000000000000	Tommotian	Padat Fm Lower Khesen -300m -= -600 Vendian Ediacaran	mineseh min	vious stratigraph
in and Tyzhinov 1998; Zhegallo et . svskaya and Letnikova, Dornamja: 2013	Ukhukapeusugal Fm theinun/Yhukapeusugal Fm Fower 1000-1350m 1000-1300m 1000-1000m 1000-1000m 1000-1000m 1000-1000m 1000-100000000000000000000	Hovsgol Cro Hovsgol Cro Hoper Khesen -2000m	oluky adat Fm -300m ⊣ -000m – -000m – -000m – -000 -000m – -000 -000m – -000 -000m – -000m – -0000m – -000m – -000m – -000m – -000m – -000m – -0000m	Darkhaf Group Preserve P	previous stratigraph
Osokin and Tyzhinov 1998; Zhegallo et. Vishnevskaya and Letnikova, Dorramja 2013	Cambrian up Ukhurologoi heinuri/Khukheusugoi Fm fomation 1000-1350m 1000-1000-1000-1000-100000000000000000	Hövsgöl Gro esen Fm E+ Upper Khesen -2000m	Vendian Vendian <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'' <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'' <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk' <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk'</sup> <sup>oluk' <sup>oluk'</sup> <sup>oluk' <sup>oluk'</sup> <sup>oluk' <sup>oluk'</sup> <sup>oluk' <sup>oluk' <sup>oluk' <sup>oluk' <sup>oluk' <sup>oluk' <sup>oluk' <sup>oluk' <sup>oluk' <sup></sup></sup></sup></sup></sup></sup></sup></sup></sup></sup></sup></sup></sup></sup></sup></sup>	Aiphean Darkhat Group Minan Mi	of previous stratigraph
Osokin and Tyzhinov 1998; Zhegallo et. Vishnevskaya and Letnikova, Dorramja 2013	Cambrian up Ukhurologi Cambrian Lower Lower 1000-1350m 1000-1000-1000000000000000000000000000	Hövsgöl Gro esen Fm Esh Upper Khesen -2000m	Vendian Vendian Vendian Vendian Vendian	Riphean Darkhat Group Oarkhat Group Minan Minan-Cryogenian Tonian-Cryogenian	m of previous stratigraph
Osokin and Tyzhinov 1998;         Zhegalio et.           2004         Vishnevskaya and Lennikova,         Dornamila:           2013         2013         2013		Hövsgöl Gro Hövsgöl Gro Hövsg	Vendian Vendian 		ation of previous stratigraph
et al, 2004 Osokin and Tyzhinov 1998; Zhegallo et. Vishnevskaya and Letnikova, Dornanija 2013	uguri series 	Khubsqull G Hövsgöl Gro 	Vendian Vendian Saar Fm Lower Khesen Saar Fm Lower Khesen Saar Fm Lower Khesen Saar Fm Lower Khesen Saar Fm Lower Khesen	анклаг series 235000 	pilation of previous stratigraph
llyin et al, 2004 Osokin and Tyzhinov 1998; Zhegallo et: Vishnevskaya and Letrnikova, Dornamja. 2013	Lambrian ugul series 	Vendian Khubsuqui G Maganian Pigerkhesen P	Vendian Vendian 100k <sup>V</sup> 100k <sup>V</sup> 100	Upper Kiphean Darkhat Series 3500m 	ompilation of previous stratigrap
Diploment         Disolation         Displayer         Zhegalio et:           Urishmerskaya and Letrnikova,         Dorramja.         2013         2013	Cambrian 	Vendian Khubsuqul G 100mmotian Tommotian	Vendian Vendian Vendian Vendian Vendian Vendian	Upper Riphean Darkhat series 3500m 35000m 3500m 3500m 3500m 3500m 3500m	Compilation of previous stratigraph
Ilyin et al. 2004 Osokin and Tyzhinov 1998; Zhegallo et: Vishnevskaya and Letnikova, Dorramja. 2013	3000m 	Vendian Khubsuqul G 100mmotian Tommotian Vendian	Pendian Vendian 100m → 600 100m → 600 100m → 600 100m → 600	- 2300m- -	3. Compilation of previous stratigraph
I 973         Display and Tyzhinov 1998;         Zhegallo et: Attnikova,         Zhegallo et: Ilyin et al. 2004         Dornanja.           181         2013         2013         2013         2013	3000m 3000m 3000m 3000m 3000m 3000m 3000m 3000m 3000m 3000m 3000m 100-1350m	Vendian Vendian Hövsgöl Gro 	Vendian Vendian Vendian Vendian Vendian	asson assonn asson assonn assonn asson assonn asson asson asson	ig. 3. Compilation of previous stratigraph
llyin, 1973 Dokin and Tyzhinov 1998; Zhegallo et: 1 and Ratnikova, Ilyin et al. 2004 Vishnevskya and Letrikova, Dornamja. 1981 2013	ugul series 	Khubsuqul G Khubsuqul G Khubs Peen Fm Peerkhesen Pommotian F Khubsuqul G F M F M Khubsuqul G	Vendian Soluki Soluk	Darkhat series 3500m	Fig. 3. Compilation of previous stratistrapl

5

transgressive sequence form the Khirvesteg The newly-defined Formation. Kheseen Formation (spelling adapted to conform to Mongolian anglicized spelling standardization) includes the Khuvsgul phosphorite and overlying carbonate grainstones and mass-flow deposits. We then define Lower, Middle, and Upper members of the Erkhelnuur Formation based on regional lithostratigraphic trends in streamline mapping, order to field identification. and structural relationships. Finally, varied siliciclastic rocks overlying the Erkhelnuur Formation are incorporated into the Ukhaatolgoi Formation.

### **METHODS**

During the course of a geological mapping campaign spanning three field seasons (summers of 2017-2019), we measured more than 8,000 m of stratigraphic section at submeter resolution through Khuvsgul Group exposures in the Khoridol Saridag Range and Darkhat Valley. Section locations are depicted in Fig. 2, while representative stratotype sections are plotted in Fig. 5. By combining our lithostratigraphic and facies observations (both from measured sections and from outcrop encountered during the course of geologic mapping) with structural relationships from our mapping efforts, we discretized constituent formations of the Khuvsgul Group into a stratigraphic framework that will streamline future work in the region. Formalized unit names, type section locations, thicknesses, age constraints, and correlations are summarized in Table 1.

### Stratigraphy

### **Ongolog** Formation

Composed of poorly sorted sandstone, siltstone, conglomerate, and diamictite, the Ongolog Formation is 250-650 m thick, with changes in both thickness and facies observed in northsouth along-strike exposures west of Lake Khuvsgul. The Ongolog Formation is well exposed on the northern slopes of Ongolog Gorge (Fig. 2), with recessive, poorly outcropping clast-free sequences of fine to medium-grained lithic arenites giving way upsection to more resistant matrix-supported stratified diamictite, with the matrix composed of the same material that makes up the clast-free interval below. On a ridge north of Kheseen Gorge (Fig. 2), a continuous section preserves the gradual transition from clast-free siliciclastic material to massive diamictite (see sections M613 and M614 in Macdonald and Jones, 2011). At both localities, the formation culminates in massive, dark-colored, argillite matrix diamcitite with clasts of carbonate, sandstone, and crystalline basement ("perforated shales" of Ilyin, 1973), depicted in in Fig. 4A.

### **Bakh** Formation

Named for its type locality on a ridgeline above Bakh Gorge (EA1701 and EA1702, Fig. 5), the Bakh (6ax; *tr. "toad"*) Formation is composed of platformal and upper slope carbonate sequences that are subdivided into three distinct members.

Khurts Member: Dolomite wackestone and finelaminated grainstone of the lowermost Khurts (хурц; *tr*. "sharp") Member of the Bakh Formation directly overly Ongolog the Formation, with a sharp, conformable or paraconformable contact. Although most Khurts Member exposures are heavily recrystallized, a generally shallowing-upward trend is suggested by the increase of coarse grainstone and grainflow deposits containing ooids and microbial mat rip-up clasts towards the top of the member. The Khurts Member is 20-120 m thick, and thickens to the southwest.

Bumbulug Member: A conformable transition to repeating shallowing-upward parasequences of limestone mudstone, marl (mixed carbonate and fine-graded siliciclastic rocks), and rhythmite (graded calcisiltite to micrite) interbeds marks the beginning of the Bumbulug (бөмбөлөг; tr. "bubbly") Member. Vitreous black-grey ellipsoidal chert nodules from 1-3 cm are distributed throughout thicker, tawny-brown mudstone beds, giving outcrops a bubbly, sponge-like almost appearance. Thinner laminated mudstone, marl, and rhythmite are more frequent up-section, with mm- to cm-scale bedded chert interspersed throughout; the Bumbulug Member is 50-150 m thick.

Name	Ongolog Formation	
Name Derivation	Good, albeit incomplete, exposures in and around Ongolog Gorge, Khuvsgul Aimag, Mongolia; previously used as the name for a broader formation that included the now-differentiated Ongolog and Bakh Formations	
Category and Rank	Lithostratigraphic Formation	
Type Area	South-facing slopes north of Kheseen Gorge, Khuvsgul Aimag, Mongoli a	
Unit Type Section	Composite section M613 and M614 (Macdonald and Jones, 2011; Fig. 2, this paper). South-facing slope near top of ridge/treeline north of Kheseen Gorge; near northeasternmost exploratory trench visible on satellite imagery. Gradational basal contact begins just west of section start (50.74887N, 100.19923E), ends just below first exposure of allodapic limestones to WNW	
Unit Description	Poorly sorted litharenites, silts, and gritty lags grade into matrix supported, beige-green to purple-weathering matrix-supported diamictite hosting sub angular to subrounded clasts (dominantly carbonates, with minor quartzite and granitic clasts) ranging from pebbles to boulders. Reported striated and faceted clasts (Osokin and Tyzhinov, 1998) suggest a glacigenic origin. Massive and stratified diamictite facies present in varying abundance throughout exposures, with rapid facies and thickness variability apparent between along-strike exposures at sub-km length scales. Uppermost strata of the Ongolog Formation include a massive diamictite with sub-rounded to sub angular carbonate, granite, and quartzite clasts supported by a dark grey argillitic matrix	
Unit Reference Sections	$1.\ M605\ (50.710609N,\ 100.178601E)\ 2.\ M610\ (50.72995N,\ 100.18197E)\ (Macdonald\ and\ Jones,\ 2011),\ located\ in\ Ongolog\ Gol\ and\ on\ ridge\ north\ of\ Ongolog\ Gol,\ respectively$	
Dimensi ons	~430m in composite type section, estimated thicknesses from ~250m in Ongolog Gorge to up to 650m to north	
Geologic Age	Earliest Cryogenian (~719Ma) to end-Sturtian (~660Ma) via correlation with globally distributed Sturtian glacial diamictites	
Regional Correlations	Maikhan-Uul Fm, Zavkhan Terrane, Mongolia (Macdonald and Jones, 2011, Bold et al., 2016a)	
Name	Bakh Formation (Khurts, Bumbulug, and Salkhitai Members)	
Name Derivation	Type section located on ridge west of Bakh Gorge, Khuvsgul Aimag, Mongolia, with best access by foot up Bakh Gol. Khurts (хурц; tr. "sharp") Member is often a cliff-forming unit; Bumbulug (бөмбөлөг; tr. "bubbly") Member contains nodular cherts that are reminiscent of bubbles on exposed weathered surfaces; Salkhitai (салхигай; tr. "windy") Member is often a ridge-forming unit, with a stratotype section on a ridgeline that is particularly exposed to wind and weather	
Category and Rank	Lithostratigraphic Formation composed of 3 component Members	
Type Area	Ridge on the westernmost rim of the upper northern drainage area of Bakh Gorge, Khuvsgul Aimag, Mongolia	
Unit Type Section	EA1701-2 (Figures 2 and 5, this paper). Lower Boundary: lowest continuous exposure of Khurts Member (50.87971N, 100.17848E) above nearby Ongolog Formation outcrop. Section follows exposures along the northern and eastern edges of ridge crests. Upper Boundary: contact with overlying yellow carbonate diamictite of Khirvesteg Formation (50.875802N, 100.179634E)	
Unit Description	Khurts Member: dolomite wackestones and laminated grainstones, often heavily recrystallized to the point of destroying primary sedimentary textures. Generally appears to shallow upsection: ooids and microbial textures more frequent towards top of member Bombolog Member: shallowing-upward parasequences of limestone, mudstone, and marl. Ellipsoidal chert nodules (1-3cm) throughout thicker mudstone beds; laminated mudstones, marls, rythmites, and bedded mm-scale cherts become frequent upsection Salkhitai Member: coarse limestone grainstone parasequences, grading into thicker dolomitized grainstones, grainflow deposits, carbonate	
Unit Reference	intraciast precess, and sedimentary precess often with lithic grains in matrix material 1. ~40m into EA1814 (50.85318N, 100.06865E), Arsai Gorge (Figures 2 and 5, this paper) 2. partial exposures in Kheseen and Ongolog Gorges (Fig. 2, this paper)	
Dimensions	At the type section, the Khurts Member is at least 20m thick, the Bombolog Member is ~120m thick, and the Salkhitai Member is ~175m thick. Elsewhere in the basin, the Khurts, Bombolog, and Salkhitai Members have been estimated to be a maximum of 120m, 150m, and 175m thick respectively, both at EA1814 and at incomplete exposures in the easternmost Khoridol-Saridag Range	
Geologic Age	Cryogenian non-glacial interlude; >660Ma to likely younger than ~657Ma (Rooney et al., 2020)	
Regional Correlations	Taishir Fm., Zavkhan Terrane, Mongolia (Macdonald and Jones, 2011; Bold et al., 2016a)	
Name	Khirvesteg Formation	
Name Derivation	Good exposure in outcrop on the south-facing hillsides west of Khirvesteg Gorge, Khuvsgul Aimag, Mongolia	
Category and Rank	Lithostratigraphic Formation	
Type Area	Ridge on the westernmost rim of the upper northern drainage area of Bakh Gorge	
Unit Type Section	~300m into EA 1701-2 (50.875683N, 100.179771E), Bakh Gorge ridge (Figures 2 and 5, this paper). Lower Boundary: locally conformable contact with underlying dolomite grainstone of Salkhitai Member, Bakh Formation (50.875802N, 100.179634E). Section continues through outcrop exposure on northern edge of ridgeline. Upper Boundary: unconformity surface at contact with Kheseen Formation (50.875516N, 100.179958E)	
Unit Description	Yellow-weathering, massive matrix-supported diamictite, containing angular to sub-angular carbonate clasts in a dolomite matrix. Often interfingered with angular dolomite boulder breccias; thickness of diamictite/breccia package highly variable in along-strike exposures. The diamictite is overlain by a 3m thick, cream-colored dolomite grainstone, often bearing sheetcrack cements and barite fans., interpreted to be a cap carbonate paired with the underlying diamictite. Above the cap carbonate, a transgressive sequence of marls, mudstones, and thinly laminated carbonaceous shales, ranging in color from white to tawny brown on weathered surfaces, is found at most exposures	
Unit Reference Sections	1. M611 (50.692094N, 100.172123E), Khirvesteg Gorge, (Macdonald and Jones, 2011). 2. M607-M609 composite section (50.691558N, 100.171421E), northern Ongolog Gorge. 3. ~200m into EA1814 (50.85510N, 100.06526E), Arsai Gorge (Figures 2 and 5, this paper)	
Dimensions	At the type section, the diamictite/dolomite breccia unit is $\sim$ 25m thick, the Marinoan cap is $\sim$ 3m thick, and the overlying Ediacaran strata are $\sim$ 20m thick. Diamictites at Reference Section 2 approach 50m in thickness, with a $\sim$ 2m thick cap carbonate sequence and nearly 50m of overlying Ediacaran strata. A round the basin, thicknesses range from 0.5 to $\sim$ 70m for the Khirvesteg diamictite, 1-3m for the Marinoan cap carbonate, and 0 to 50m for the Ediacaran strata	
Geologic Age Regional Correlations	Onset of Marinoan Glaciation (younger than 657Ma; Rooney et al., 2020) to early Ediacaran (younger than 635Ma) Khongor Formation and lower Ol Formations, Zavkhan Terrane, Mongolia (Macdonald and Jones, 2011; Bold et al., 2016a)	

### Table 1. Formalization of newly defined stratigraphic units of the Khuvsgul Group

7

Name	Kheseen Formation
Name Derivation	Good exposure in and around Kheseen Gorge, Khuvsgul Aimag, Mongolia; previously used as the name for a broader formation (frequently spelled "Khesen") containing the now-differentiated Khirvesteg Formation and the new Kheseen Formation
Category and Rank	Lithostratigraphic Formation
Type Area	Ridge on the westernmost rim of the upper northern drainage area of Bakh Gorge
Unit Type Section	~345m into EA1701-2, (50.875516N, 100.179958E) Bakh Gol Ridge (Figures 2 and 5, this paper). Lower Boundary: carbonates directly above unconformity surface at contact with Kheseen Formation (50.875516N, 100.179958E). Section continues along fall-line exposure directly up ridge to east. Upper Boundary: end of continuous exposure before contact with overlying Erkhelnuur Formation
Unit Description	Thin, interbedded limestone and dolomite mudstones, fetid marls, and phosphatic and silicified hardgrounds form discontinuous beds frequently truncated by cm-scale channelization and autoclastic debris flows. These lower Kheseen Formation strata contain abundant primary phosphate precipitates, and, in addition to hosting a DPT-type microfossil assemblage (Anderson et al., 2017,2019), also contain stromatolitic and microbial mat textures. Thin interbeds of the lower Kheseen Formation are usually overlain by a marker bed of amorphous chert, which is subsequently overlain by allodapic carbonate grainstones bearing phosphatic and siliceous allochems. While exposure at the type section locality ends within the allodapic carbonate sequence, exposures at Reference Section 1 reveal an expanded carbonate grainstone sequence in which phosphatic and siliceous allochems become sparse upsection before the occurrence of a second thick bed of amorphous chert. At Reference Section 2, the uppermost allodapic carbonates begin to exhibit a strong siliciclastic allochemical component, eventually giving way to graded lithic grainflows and larger debris flows with up to boulder-sized clasts. Here, the contact with the overlying Erkhelnuur Formation is a microkarstic disconformable exposure surface
Unit Reference Sections	1. ~ 230m  into EA1814 (50.855144N, 100.065221E), Arsai Gorge (Figures 2 and 5, this paper); 2. EA1905 (50.855144N, 100.18400E), Kheseen Gorge (Figures 2 and 5, this paper)
Dimensions	At least 130m thick at the type locality; at least 500m thick at Reference Section 1, and between 120 and 400m at other exposures around the basin
Geologic Age	Latest Ediacaran to earliest Cambrian via correlation with Zavkhan equivalents (Macdonald et al, 2009; Macdonald and Jones, 2011)
Regional Correlations	Zuun Arts Formation, Zavkhan Terrane, Mongolia (Macdonald and Jones, 2011; Smith et al., 2016)
Name	Lower, Middle and Upper Members of the Erkhelnuur Formation
Name Derivation Category and Rank	Relative stratigraphic heights of differentiable units of the Erkhelnuur Formation 3 constituent Members of a Lithostratigraphic Formation
Type Area	Lower Member: ridges west of Khirvesteg Gorge; Middle Member: ridges west of Khirvesteg and Ongolog Gorges; Upper Member: ridges west of Ongolog and Kheseen Gorges, as well as in several locations in the upper Arsai Gorge drainage
Unit Type Section	Lower Member: EA1801-2 (50.70571N, 100.12742E), ridge west of Khirvesteg Gol (Figures 2 and 5, this paper). Lower Boundary: base of EA1801-2, located ~20 m above contact with Kheseen Formation, which is covered by vegetation. Section follows ridgecrest to North. Upper Boundary: ~280m into EA1801-2 Middle Member: begins at upper boundary of Lower Member. Section continues north, following ridge crest, before turning east at (50.723044N, 100.136873E) and continuing along the ridgeline to the east, into EA1825, which follows the same ridge. Upper Boundary: ~650m into EA1825 (50.724445N, 100.1366252E), ridge west of Ongolog and Kheseen Gorges (Figures 2 and 5, this paper).
	Upper Member: begins at upper boundary of Middle Member. Exposure continues until EA1825 ends at core of syncline. Uppermost Upper Member is exposed in EA1920 (50.76544N, 100.03005E), Upper Arsai Gorge drainage (Figures 2 and 5, this paper). Section continues up gully until conformable contact with overlying Ukhaatolgoi Formation Lower Member: bounded below by a disconformable exposure surface at the top of the Kheseen Formation, the Lower Member of the Erkbelmur Formation is composed of dolomite and dolomitized limestone shallowing-upward paragemences (typically
Unit Description	laminated mudstones and grainstones, coarser grainstones with infrequent stromatolitic horizons, and allodapic packstones and coarse grainstones containing ooids and oncoids. Occasional silicified surfaces Middle Member: abrupt transition to parasequences dominated by dark-grey-weathering limestone grainstone beds; this gray band is easy to identify on satellite or aerial imagery. Ichnofossils begin to appear, with abundant bioturbation increasing up section, eventually obfuscating all primary sedimentary structures and bedding features. Archaeocyathid reefs and hash appear ~300m above the base of the Middle Member Upper Member: base of the member is demarcated by an abrupt transition to white-cream dolomite grainstone beds. Heavy recrystallization throughout, with occasionally discernible laminated grainstones and rhythmites. White dolomites are also visible in satellite and aerial imagery, and serve as a valuable marker bed. The uppermost portion of the Upper Member begins to incorporate lithic grains and fragments in grainstone beds, with the frequency of lithics increasing upsection until carbonate deposition is overwhelmed at the conformable contact with the overlying Ukhaatolgoi Formation
Unit Reference Sections	1. ~119m into EAGC1905 (50.741915N, 100.183101E), Kheseen Gol (Figures 2 and 5, this paper), which includes the disconformable contact at the base of the Lower Member of the Erkhelmur Formation
Dimensions Geologic Age Regional Correlations	Lower Member: ~300m thick; Middle Member: 600-800m thick; Upper Member: 500-600m thick Archaeocyathid appearances constrain the Middle Member to within the Tommotian and Toyonian Stages (~525-511Ma) Putative correlation with Bayangol Formation, Zavkhan Terrane, Mongolia (Smith et al., 2016)

 Table 1. Formalization of newly defined stratigraphic units of the Khuvsgul Group (continue)

Salkhitai Member: The uppermost portion of the Bakh Formation, the Salkhitai (салхитай; tr. "windy") Member, conformably overlies the Bumbulug Member and contains coarse limestone grainstone-dominated parasequences that transition into a coarsening-upward stack of dolomitized grainstone, grainflow deposits, and intraclast breccia with a minor siliciclastic granular component. The top of the Salkhitai Member is marked by a coarse, carbonatecemented sandstone with carbonate granules that is overlain by a massive unit of dolomite grainstone. The Salkhitai Member ranges between 100 m and 150 m in thickness.

#### Shar Formation

The Shar (map; tr. "yellow") Formation is composed of a massive matrix-supported



**Fig. 4.** Khuvsgul Group rocks in outcrop and hand-sample. **A:** Matrix-supported diamictite of the Ongolog Formation, featuring sub-angular to sub-rounded carbonate and granitic clasts. **B:** Yellow, dolomite-matrix-supported diamictite of the Shar Formation, with angular to sub-angular carbonate clasts, many of which are lithologically similar to strata of the underlying Bakh Formation. **C:** phosphorite of the lower phosphatic units of the Kheseen Formation observed in EA1701, with discontinuous layers of primary phosphatic material weathering blue-white between darker chert and limestone beds. **D:** Dolomite grainstones of the Kheseen Formation bearing phosphatic allochems, EA1814. **E:** Bed-penetrating bioturbation within grainstones of the Middle Member of the Erkhelnuur Formation. **F:** Archaeocyathid hash in grainstones of the Middle Member of the Erkhelnuur Formation.

diamictite with a distinctive yellow-ochre dolomite matrix (Fig. 4B), often closely associated with coarse, gray-weathering angular dolomite breccia. Around the basin, thickness of diamictite-breccia packages ranges from 0.5-70 m. The base of the member is defined as the transition from dolomite grainstone of the uppermost Bakh Formation to diamictite or dolomite breccia. At the stratotype locality above Bakh Gorge, the Shar Formation is approximately 20 m thick, with nearly ten meters of angular dolomite breccia overlain by the eponymous yellow diamictite (EA1701-2, Fig. 5).

### Khirvesteg Formation

Conformably above the diamictite and breccia, the base of the Khirvesteg Formation is defined by a 1-3 m thick dolomite grainstone containing barite fans and sheetcrack cements (Hoffman et al., 2011). The cap carbonate is overlain by thin dolomitic marl, mudstone, and carbonaceous shale. Although exposures to the south are locally only 2-3 m thick, 25 m of marl, shale, and dolomite mudstone are found in the stratotype section on the ridge above Bakh Gorge (EA1701-2, Fig. 5). Note that in Fig. 2, the Shar and Khirvesteg Formations are shown as a single map unit for visual clarity at the presented map scale.

### Kheseen Formation

Ranging from 160-500 m in thickness, the Kheseen Formation is separated from the Khirvesteg Formation by a prominent erosional unconformity. Above this surface, the basal Kheseen Formation is composed of interbedded black micritic limestone and dolomite mudstone, foetid marl and shale, and phosphatic and silicified hardgrounds (Fig. 4C). Bedding is largely discontinuous along strike, and is often truncated by channelization or autoclastic debris flows, while thicker carbonate beds contain stromatolitic thrombolitic and textures. Phosphatic hardgrounds and phosphatic grainflows in these strata contain a Doushantuo-Pertatanka-Type microfossil assemblage, described by Anderson et al. (2017; 2019). This sequence, referred to in previous works as the lower phosphate unit (Ilyin and Ratnikova,

1981; Ilvin, 1998), is 10-60 m thick, and contains the most concentrated phosphate ore (average of 22% and up to 34% P<sub>2</sub>O<sub>5</sub>; Dorinamiaa and Altanshagai, 2015) in the region. This phosphatic unit is capped by 1-5 m of massive amorphous black chert. Above this, granular phosphorite is present in graded allodapic limestone and dolomite grainstone (Fig. 4D). This facies forms the majority of phosphorite exposure in the Khuvsgul Basin. In the Western Khoridol Saridag Range (KSR), the Kheseen Formation is capped by a second chert bed, whereas in the easternmost KSR it is composed of grainstone and coarse conglomerate with a large siliciclastic component not observed elsewhere in the basin. The stratotype section for the Upper Kheseen Formation is on the ridge above Bakh Gorge (EA1701-2) with reference sections in Arsai Gorge (EA1814), which includes an expanded upper phosphorite sequence, and in Ongolog Gorge (EA1905), which contains conglomerate with abundant lithic clasts at the top of the formation. All sections are depicted in Fig 5.

### Erkhelnuur Formation

The Erkhelnuur Formation (Zhegallo, 2000) is a 2 km thick carbonate sequence separated from the Kheseen Formation by a hiatal surface, observed in the Eastern KSR (EA1905) as a disconformable exposure surface, with the base of the Erkhelnuur Formation defined as the carbonate grainstones overlying this surface. In order to simplify structural interpretation and lithological differentiation of Erkhelnuur Formation carbonates in the field, we propose the subdivision of the formation into Lower, Middle, and Upper Members, which are summarized below.

*Lower Member:* The Lower Member includes repetitive dolomite and dolomitized limestone parasequences, composed of laminated mudstone and grainstone interbeds, dolomitized grainstone beds containing domal stromatolites, and allodapic packstone and coarse grainstone containing ooids and oncoids. Parasequences are commonly capped by horizons of lenticular grey -black chert. The Lower Member is 250-300 m



**Fig. 5.** Measured stratigraphic sections from Khuvsgul Group type-localities. All section locations are depicted in Fig. 2, while coordinate locations shown above are from the stratigraphic base of each section.

thick, with a stratotype exposure on a ridge above Khirvesteg Gorge (EA1801-2, Fig. 5).

Middle Member: The base of the Middle Member of the Erkhelnuur Formation is defined by an abrupt transition to parasequences dominated by dark grey limestone grainstone the parasequences beds. While include stromatolitic mudstone and dolo-rhythmite, this transition is visible both in the field and on satellite imagery, where limestone-dominated parasequences appear as a dark blue-gray band. In most sections, bed-penetrating ichnofossils appear 20-50 m above the base of the Middle Member, consisting of irregular 1-2 cm diameter tubes. The density of bioturbation increases up-section to the point of obfuscating (Fig. primary bedding features 4E). Approximately 300 m above the base of the Middle Member, archaeocyathid reefs and hash can be observed in bioturbation-free zones (Fig. 4F), constraining the strata to a window of the early Cambrian between the Tommotian and Toyonian Stages (~525-511 Ma, or mid-Terreneuvian Age 2 to Cambrian Series 2, Age 4; Peng et al., 2020). The Middle Member of the Erkhelnuur Formation ranges from 600-800 m thick, and is best exposed in sections EA1801-1802 (see Fig. 5) on the ridges above Khirvesteg Gorge.

Upper Member: The base of the Upper Member of the Erkhelnuur Formation is demarcated by a sequence of >50 m thick white- to creamcolored dolomite grainstone beds. Although the white beds locally preserve laminated grainstone and rhythmite, in most exposures, primary bedding features are obliterated by recrystallization. Like the dark beds at the base of the Middle Member, the basal white dolomite of the Upper Member forms a reliable marker bed that is easily visible in the field and satellite imagery. Above the white dolomite, grey dolomite and limestone laminite-grainstonegrainflow parasequences continue, with ichnofossils present in fine grainstone beds. Upsection, coarser-grained grainstone and grainflows incorporate terrigenous lithic fragments, which increase in frequency and size toward the top of the section. The Upper

Member is 500-600 m thick; the base of the Member is best exposed in EA1825, while the uppermost grainflows are prominently exposed in section EA1820 (Fig. 5).

### Ukhaatolgoi Formation

The >350 m thick Ukhaatolgoi Formation conformably overlies the Erkhelnuur Formation and is composed of siliciclastic rocks dominated by greywacke with minor siltstone, sandstone, and conglomerate. Carbonate parasequences of the uppermost Erkhelnuur Formation contain allochems that are gradationally succeeded by greywacke of the basal Ukhaatolgoi Formation. Previously referred to as the Ukhutologov Formation (Zhegallo et al., 2000; Demidenko et al., 2003) and excluded from the Khuvsgul Group, the Ukhaatolgoi Formation is dominated by green greywacke, with rare granule to pebble lithic clasts, angular quartz and plagioclase grains, and carbonate fragments in a green siltstone matrix. The base of the formation is best documented in section EA1820, and the formation is best exposed at the cores of synclines in upper Arsai Gorge. Note that we do not include the formalization of this Formation in Table 1, as this nomenclature has been applied to these rocks by previous workers (Dorjnamjaa et al., 2015).

### **DISCUSSION**

While the stratigraphic divisions outlined above are largely based on lithological characteristics that can be easily differentiated in the field by future workers, our stratigraphic framework is also reflective of significant changes in paleoenvironment and paleoclimate that occurred during the deposition of the Khuvsgul Group. For example, the Ongolog, Bakh, Shar, and Khirvesteg formations can be correlated with the Sturtian glaciation, Cryogenian nonglacial interlude, Marinoan Glaciation, and basal Ediacaran transgression respectively, with diamictites associated with both glaciations separated by a platformal carbonate sequence (Macdonald and Jones, 2011). A glaciogenic origin for the Ongolog Formatiton is established via the observation of exotic clasts in matrixsupported diamictite, as well as striated and faceted clasts (Osokin and Tyzhinov, 1998). Typically exhibiting sheetcrack cements and crystal fans, the basal carbonate of the Khirvesteg Formation is interpreted as a cap carbonate sequence (Hoffman et al., 2011), coupled with the underlying diamictite of the Shar Formation. Additionally, these units can be correlated with Cryogenian and Ediacaran successions on the Zavkhan Terrane in southwest Mongolia (Macdonald and Jones, 2011; Bold et al., 2016b).

Bound above and below by unconformities, the newly-defined Kheseen Formation records a depositional environment that resulted in the precipitation and eventual reworking of the Kheseen phosphorites, which may be condensed equivalents of the phosphatic shales of the Zuun Arts Formation on the Zavkhan Terrane (Macdonald and Jones, 2011). The lower units of the Kheseen Formation record precipitation of primary phosphatic material in a very shallow, energetic depozone, whereas phosphatic material in the upper parts of the formation has been redeposited in carbonate grainstones (Fig. 5). Importantly, the newly defined lower bound of the Kheseen Formation isolates the phosphatic strata from the underlying Khirvesteg and Shar formations, highlighting the lack of a previously-postulated genetic relationship between the Marinoan glaciation and phosphogenesis (Osokin and 1998). Tyzhinov, Furthermore, the disconformable exposure surface at the top of the formation separates the Kheseen Formation in time from the nearly 2 km of overlying Erkhelnuur Formation carbonates, introducing the possibility that the Kheseen and Erkhelnuur formations were accommodated by distinct basin-forming events. Additional petrographic, geochronological and chemostratigraphic data are necessary to better constrain the ages and depositional environments of the Kheseen phosphorites and Khuvsgul Group strata as a whole. The stratigraphic framework presented here provides a necessary foundation for future work in the Khuvsgul region.

### CONCLUSIONS

A new stratigraphic framework for the Khuvsgul Group divides strata spanning nearly 200 My of Earth history into seven formations. Cryogenian strata are divided into the Ongolog, Bakh. Shar. and Khirvesteg formations. discretizing units associated with the Sturtian glaciation, Cryogenian non-glacial interlude, Marinoan Glaciation, and basal Ediacaran transgression respectively. Phosphatic strata, bounded above and below by depositional hiatuses, are included in the newly defined Kheseen Formation. In order to streamline mapping and correlation, the Erkhelnuur Formation is divided into Lower, Middle, and Upper Members, while the overlying siliciclastic sequences are discretized as the Ukhaatolgoi Formation.

### ACKNOWLEDGEMENTS

Thanks to Erdene Bayarsaikhan, Ekv Erdene, Sam Lobianco, Peter Otness, and Judy Pu for assistance, companionship, and camaraderie in the field, to M. Munkhbaatar, Batsukh Erdene, and Set for transportation and friendship, to Ariunsanaa Dorj and the Earth Science Center of Mongolia for logistic help, and to the Ministry of Environment of Mongolia and the rangers and staff of the Khuvsgul Nuur National Park and Khoridol Saridag Protected Zone for access. This work was funded by grant funds from Nasa Astrobiology: Exobiology and Evolutionary Biology, NNH10ZDA001N-EXO and the NSF-GRFP.

### REFERENCES

Anderson, R.P., Macdonald, F.A., Jones, D.S., McMahon, S., Briggs, D.E. 2017. Doushantuo -type microfossils from latest Ediacaran phosphorites of northern Mongolia: Geology, v. 45(12), p. 1079-1082.

https://doi.org/10.1130/G39576.1

- Anderson, R.P., McMahon, S., Macdonald, F.A., Jones, Briggs, D.E. D.S. 2019. Palaeobiology of latest Ediacaran phosphorites from the upper Khesen Formation, Khuvsgul Group, northern Mongolia: Journal of Systematic Palaeontology, v. 17(6), p. 501-532. https://doi.org/10.1080/14772019.2018.1443977
- Bold, U., Smith, E.F., Rooney, A.D., Bowring, S.A., Buchwaldt, R., Dudás, F.Ő., Ramezani, J., Crowley, J.L., Schrag, D.P., Macdonald, F.A. 2016a. Neoproterozoic stratigraphy of

the Zavkhan terrane of Mongolia: The backbone for Cryogenian and early Ediacaran chemostratigraphic records: American Journal of Science, v. 316(1), p. 1-63.

https://doi.org/10.2475/01.2016.01

- Bold, U., Crowley, J.L., Smith, E.F., Sambuu, O., Macdonald, F.A. 2016b. Neoproterozoic to early Paleozoic tectonic evolution of the Zavkhan terrane of Mongolia: Implications for continental growth in the Central Asian orogenic belt: Lithosphere, v. 8(6), p. 729-750. https://doi.org/10.1130/L549.1
- Bold, U., Isozaki, Y., Aoki, S., Sakata, S., Ishikawa, A., Sawaki, Y., Sawada, H. 2019. Precambrian basement, provenance implication, and tectonic evolution of the Gargan block of the Tuva-Mongolia terranes, Central Asian Orogenic Belt: Gondwana Research, v. 75, p. 172-183.

https://doi.org/10.1016/j.gr.2019.05.003

- Demidenko, Y.E., Zhegallo, E.A., Parkhaev, P.Y., Shuvalova, Y.V. 2003. Age of phosphorites from the Khubsugul basin (Mongolia): Dolkady Earth Sciences C/C of Dolkady Akademiia Nauk, v. 389, p. 317-321.
- Altanshagai, Dorjnamjaa, D., G. 2015. original viewpoint Concerning the of biogeologic accumulation of the old bedded phosphorites in the Khubsugul and Zavkhan basins of Mongolia: Open Journal of Geology, v. 5(09), p 666.

https://doi.org/10.4236/ojg.2015.59059

- Dorjnamjaa, D., Enkhbaatar, B., Altanshagai, G. 2015. Concerning Precambrian and Cambrian regional stratigraphy of Mongolia: Mongolian Geoscientist, v. 42, p.2-11.
- Hoffman, P.F., Macdonald, F.A., Halverson, G.P. 2011. Chemical sediments associated with Neoproterozoic glaciation: iron formation. cap carbonate, barite and phosphorite: Geological Society, London, Memoirs, v. 36(1), p.67-80. https://doi.org/10.1144/M36.5
- Ilvin, A.V. 1971. About the Tuva-Mongolian massif: Materials about the Regional Geology of Africa and Asia: Zarubezhgeologiya, v. 22, p.67-73.
- Ilvin, A.V. 1973. Khubsugul phosphoritebearing basin: Moscow, Geolgicheskiy Institut, Doklady Akademiya Nauk SSSR, p. 167 (in Russian).

- Ilvin, A.V. 1998. Rare-earth geochemistry of old phosphorites and probability of syngenetic precipitation and accumulation of phosphate: Chemical Geology, v. 144(3-4), p. 243-256. https://doi.org/10.1016/S0009-2541(97)00134-4
- Ilyin, A.V., Ratnikova, G.I. 1981. Primary, bedded, structureless phosphorite of the Khubsugul Basin, Mongolia: Journal of Sedimentary Research, v. 51(4), p. 1215https://doi.org/10.1306/212F7E69-2B24-1222. 11D7-8648000102C1865D
- Kuzmichev, A.B., Bibikova, E.V., Zhuravlev, D.Z. 2001. Neoproterozoic ( $\square$  800 Ma) orogeny in the Tuva-Mongolia Massif (Siberia): island arc-continent collision at the northeast Rodinia mar-gin: Precambrian Research, v. 110(1-4), p. 109-126. https://doi.org/10.1016/S0301-9268(01)00183-8

2015. Kuzmichev, Neoproterozoic A.B. accretion of the Tuva-Mongolian massif, one of the Precambrian terranes in the Central Asian Orogenic Belt: The Central Asian Orogenic Belt, p. 66-92.

Macdonald, F.A., Jones, D.S. 2011. The Khubsugul Group, northern Mongolia: Geological Society, London, Memoirs, v. 36 (1), p. 339-345.

https://doi.org/10.1144/M36.30

- 1998. Osokin, P.V., Tyzhinov, A.V. Precambrian tilloids of the Oka-Khubsugul phosphorite-bearing basin (eastern Sayany, northwestern Mongolia): Lithology and Mineral Resources, C/C of Litologiia I Polezvne Iskopaemye, v. 33, p. 142-154.
- Peng, S.C., Babcock, L.E., Ahlberg, P. 2020. The Cambrian period in Geologic Time Scale 2020. Elsevier. 565-629. https:// p. doi.org/10.1016/B978-0-12-824360-2.00019-X
- Smith, E.F., Macdonald, F.A., Petach, T.A., Bold, U., Schrag, D.P. 2016. Integrated stratigraphic, geochemical, and paleontological late Ediacaran to early Cambrian from records southwestern Mongolia: GSA Bulletin, v. 128(3-4), p.442-468. https://doi.org/10.1130/B31248.1
- Vishnevskaya, I.A., Letnikova, E.F. 2013. Chemostratigraphy of the Vendian-Cambrian carbonate sedimentary cover of the Tuva-Mongolian microcontinent: Russian Geology and Geophysics, v. 54(6), p. 567-586. https://doi.org/10.1016/j.rgg.2013.04.008

Zhegallo, E.A., Rozanov, A.Y., Ushatinskaya G.T., Hoover, R.B., Gerasimenko, L.M., Ragozina, A.L. 2000. Atlas of microorganisms from ancient phosphorites of Khubsugul (Mongolia): Huntsville, AL, National Aeronautics and Space Administration, Marshall Space Flight Center, 166 p.